

Prepared by the Army Map Service (AMS), Corps of Engineers, U.S. Army, Washington, D.C. Compiled in 1954 by photogrammetric methods. Horizontal and vertical control by USCGS and USGS. Aerial photography 1954. Photography field annotated 1954.

REFERENCES

Briggs, I. C., 1974. Machine contouring using minimum curvature. *Geophysics*, v. 39, p. 39-48.

Hammer, S., 1939. Terrain corrections for gravimeter stations. *Geophysics*, v. 4, p. 184-194.

Morelli, C., 1976. Modern standards for gravity surveys. *Geophysics*, v. 41, p. 1051.

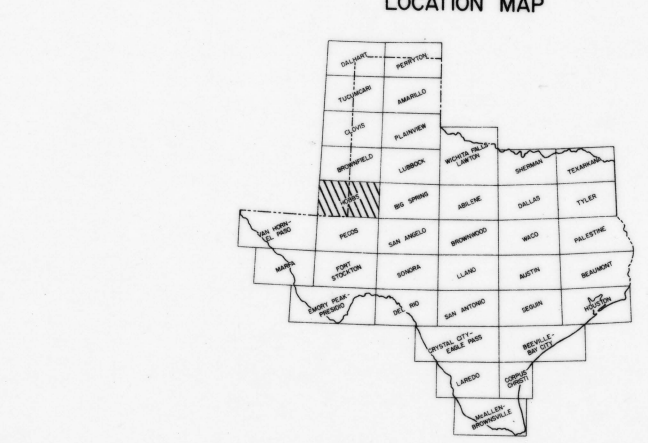
Plouff, D., 1977. Preliminary documentation for a FORTRAN program to compute gravity terrain corrections based on topography digitized on a geographic grid. U.S. Geological Survey, Open-File Report 77-538, 45 p.

Sampson, R. J., 1978. Surface II graphics system. Kansas Geological Survey, 240 p.

10,000-meter Universal Transverse Mercator grid ticks, zone 13
+ = approximate location of station points

LOCATION MAP

Major Sources of Data for the Hobbs Sheet
Defense Mapping Agency; The University of Texas at Dallas; The University of Texas at El Paso; Texas Tech University; Mr. Hart Brown; Shurbet, D. H., 1966. Gravity field and isostatic equilibrium of the Llano Estacado of Texas and New Mexico. *Geol. Soc. Am. Bull.*, v. 77, p. 213-222.



BOUGUER GRAVITY ATLAS OF TEXAS, HOBBS SHEET

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TEXAS GRAVITY DATA BASE AND REDUCTION PARAMETERS

The data base of gravity readings that made the Texas gravity mapping project possible is the result of the efforts of many individuals and groups; ultimately it will contain more than 50,000 readings. Major sources of data for each sheet are indicated in the explanation. However, a particular note of thanks is due to Mr. and Mrs. Hart Brown because the regional data they so generously provided are the foundation of this project. Without these data, the mapping and compilation required would have been impossible.

Data in the Texas gravity data base are carefully edited and maintained in a standard format and are tied to a common gravity datum (IGSN-71; Morelli, 1976). Sea level was used as an elevation datum, and a density of 2.67 g/cm³ was used in the Bouguer correction. For stations west of 103°W longitude, outer-zone terrain corrections were calculated for zones extending from 0.89 km to 167 km from each station (Hammer, 1939). These calculations employed a terrain correction program written by Plouff (1977) and a set of averaged elevations on a grid interval of 30 seconds of latitude and longitude.

Thus, the data base is internally consistent, and the resulting maps all tie together at common boundaries. These maps were constructed by gridding the data using the minimum curvature technique (Briggs, 1974) and were contoured by using a modified version of the Surface II graphics system (Sampson, 1978).

ρ = Ratio of reduction density to the standard value of 2.67 g/cm³

G_{obs} = Observed gravity, in milligals, relative to the IGSN-71 gravity datum (Morelli, 1976)

G_{th} = Theoretical gravity on the surface (sea level) of the 1967 reference spheroid. $G_{th} = 978031.843 + \sigma + (1572.66 + \sigma) \sin^2 \phi + (6083.534 + \sigma) \sin^4 \phi - (1089.748 + \sigma) \sin^6 \phi$, where $\sigma = 0.0001 \rho^2$, ϕ = latitude in degrees (International Association of Geodesy, 1967; see Morelli, 1976)

C_{th} = Free air correction for the elevation of the station relative to sea level. $C_{th} = h(0.30877 + \rho(-0.001398 + \rho(0.001353 + \rho(-0.0005329 + \rho(0.000091))))$, where h = elevation of station in meters

C_{slab} = Bouguer correction for rock mass positioned between the station and sea level. Correction is calculated from the formula for the attraction of an infinite horizontal slab of thickness h and density ρ to obtain the equation $C_{slab} = -0.119h$

C_{curv} = Curvature correction, a modification of the Bouguer slab approximation that corrects for the curvature of the Earth's surface. $C_{curv} = h(1.4689108 \times 10^{-7} + h(-3.53275 \times 10^{-7} + h(4.46646 \times 10^{-7})))$

C_{rel} = Correction for local topography

All gravity data were reduced to Bouguer anomaly values using the following formulas:

$BA = G_{obs} - G_{th} + C_{th} + (C_{slab} + C_{curv} + C_{rel}) \rho$

where
BA = Bouguer anomaly